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Morphological and geochemical features of crater lakes in Costa Rica: an overview

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ABSTRACT

This paper describes the compositional and morphological features of the crater lakes found in the volcanoes of Rincón de La Vieja, Poás, Irazú, Congo and Tenorio volcanoes (Costa Rica). As evidenced by the distribution of the water and dissolved gas chemistry along vertical profiles, the different fluid sources feeding the lakes reflect the present status of each of the volcanic systems. The chemical features of the Caliente (Poás volcano) and Rincón crater (Rincón de la Vieja volcano) lakes are mainly dependent on i) inputs of magmatic fluids from sub-lacustrine fumaroles and ii) water-rock interaction processes. Conversely, the Irazú lake is mainly affected by the presence of CO₂(H₂S)-rich fluids discharged from a hydrothermal system, which masked possible magmatic fluid contributions. Rainfall and organic activity are the main factors responsible for the chemical composition of Hule, Botos, Congo and Tenorio lakes. The chemical and isotopic water composition of Botos, Irazú and Hule lakes have displayed no significant variations along the vertical profiles. In contrast, Caliente lake shows a distinctive chemical stratification, mainly involving F, Cl and SO₄²⁻. The behaviour of these compounds seems to be governed by both dissolution of highly acidic species, i.e. HF, HCl and SO₂ released from the magmatic environment, and microbial activity. Despite the significant increases with depth of dissolved CO₂ at Caliente and Irazú lakes, the hazard for Nyos-type gas eruptions can be considered negligible, since i) the water volumes are too small and ii) the convective heat transfer limits the CO₂ recharge rate. The relatively high concentrations of dissolved CO₂ measured at the maximum depth of the Hule lake are likely produced by both degradation of organic material and degassing from a deep source. The sporadic episodes of fish deaths recently observed in this lake can be associated with lake overturn processes that have favoured the rise up to the lake surface of deep, oxygen-depleted waters.

Key words: lake stratification, limnic eruptions, hyperacidic lake

1. INTRODUCTION

Crater lakes represent unique natural systems at the Earth's surface. They are commonly present in volcanic systems characterized by recent activity, being found in more than 12% of the 714 post-Holocene volcanoes of the world (Simkin *et al.* 1981). Crater lakes increase volcanic hazard because uprising magmas may directly interact with lake water, leading to high-energy phreato-magmatic eruptions, e.g. Taal volcano, Philippines (Blong 1984). Moreover, the ejected volcanic products displace the lake waters, enhancing the effects of secondary phenomena, such as the lahars, e.g. Mt Ruapehu, New Zealand (Nairn *et al.* 1979) and Mt Kelut, Indonesia, (Badrudin 1994). Crater lakes hosted within quiescent volcanic systems interact with magmatic-hydrothermal fluids, acting as condensers and calorimeters for acidic volatiles and heat from magmatic degassing (e.g., Brantley *et al.* 1993; Anzidei *et al.* 2008). This implies that contributions from sub-lacustrine gas vents, coupled with bacterial activity, may produce large CO₂(CH₄)-rich dissolved gas reservoirs (Martini *et al.* 1994; Kusakabe 1996; Schmid *et al.* 2005). Destabilization of stratified layers at depth may cause turnover of

the lake (Rice 2000) induced by either heat flux through the system or landslides from the inner crater flanks triggered by heavy rains or seismic events. At the surface, lake turnover produces the so-called "limnic eruptions" (Kusakabe 1996 and references therein), such as those occurred at Lakes Monoun and Nyos (Cameroon) in 1984 and 1986, respectively (Sigurdsson *et al.* 1987; Evans *et al.* 1993, 1994). Volcanic lakes affected by hydrothermal /magmatic fluid discharges are frequently dynamic hydrologic systems, and the high secondary permeability of the volcanic sequences greatly enhances fluid circulation (Sanford *et al.* 1995; Varekamp *et al.* 2000). Consequently, the physical-chemical characteristics of crater lakes are often useful to monitor and forecast volcanic events in active and dormant systems (Giggenbach 1974, 1983; Rowe *et al.* 1992; Christenson 2000; Anzidei *et al.* 2008).

Costa Rica is a land of volcanoes, eight of which are characterized by the presence of crater lakes (Haberyan *et al.* 2003). There are lakes within i) dormant volcanoes (Congo and Barva volcanoes), ii) systems showing moderate hydrothermal activity (Irazú and Tenorio volcanoes) and iii) craters with strong fumarolic emissions and where periodic phreatic eruptions have recently occurred (Rincón de la Vieja and Poás volcanoes).

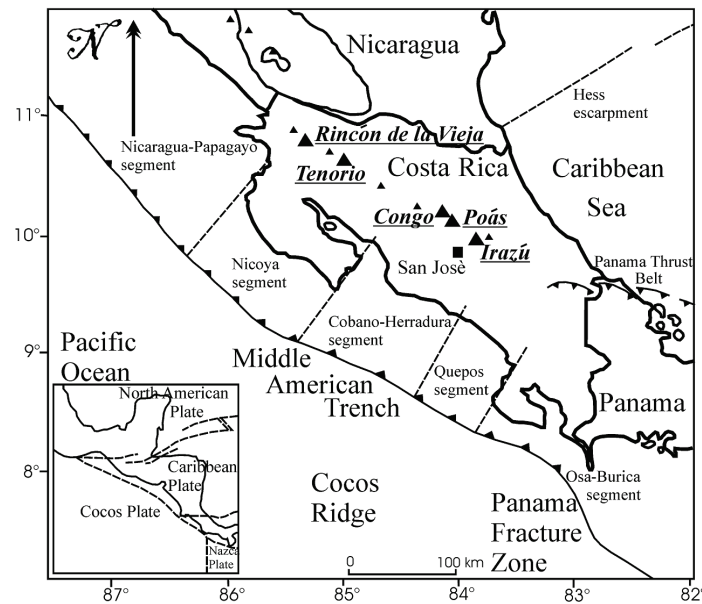


Fig. 1. Schematic map of Costa Rica and location of the main crater lakes.

This paper presents the main morphometric parameters and vertical distribution of the chemical and isotopic composition of water and dissolved gases at the Hule and Congo (Bosque Alegre caldera, Congo volcano), Irazú (Irazú volcano), Botos and Caliente (Poás volcano), Tenorio (Tenorio volcano) and Rincón Crater (Rincón de la Vieja volcano) lakes. The aim is to constrain the main processes controlling the physical-chemical features of these basins and their possible relations with the activity level of the hosting volcanic systems, by using previously published (Vaselli *et al.* 2003a; Tassi *et al.* 2005) and original data.

2. GEOLOGICAL SETTING AND VOLCANIC HISTORY

Costa Rica is characterized by a complex tectonic setting related to the subduction of the Cocos Plate beneath the Caribbean Plate (Fig. 1) at a rate of about 10 cm y^{-1} (Johnston & Thorkelstone 1997). This collisional process has caused, since the Middle Tertiary, an intense volcanic activity resulting in the formation of a Quaternary volcanic range. This includes i) the Cordillera de Guanacaste to the NW and ii) the Cordillera Central to the SE, both lying on a regional ignimbritic plateau (Montero *et al.* 1995). The former consists of five stratovolcanoes (Cacao-Orosi, Rincón de la Vieja, Miravalles, Tenorio and Arenal) mainly composed of basaltic andesite to andesitic products. The Cordillera Central has five stratovolcanoes (Platanar-Porvenir, Poás, Barba, Irazú, and Turrialba) with lava flows and pyroclastic deposits, whose composition varies from basaltic to dacitic and andesitic (Castillo 1984). The volcanic activity in Costa Rica in the last 25,000 years has mainly been characterized by phreatic and phreatomagmatic eruptions, with rare plinian eruptions (Alvarado *et al.* 1992).

2.1. Poás volcano

Poás (2708 m a.s.l.) is part of the Cordillera Central, 35 km NE of San José (Fig. 1), and is one of the most active volcanoes of Costa Rica. The summit area comprises three N-S oriented cones: 1) Von Franzius to the N, 2) Active Crater, which contains an extremely acidic lake (Caliente lake), and 3) Botos to the SE, where a second lake (Botos) is hosted. Poás volcanic edifice is mainly consisting of calc-alkaline basaltic and andesitic volcanic products (Prosser & Carr 1987). The last lava flow (7540 years B.P.) was produced by the Botos cone (Prosser 1985). Poás historical activity (since 1828) has been characterized by geyser-like and rare phreatomagmatic eruptions from the Active Crater (Krushensky & Escalante 1967; Casertano *et al.* 1983). The material ejected during the 1953-54 explosive events formed a 40-m-high cone that presently occupies most of the NW sector of the Active Crater. In April and May 1989 and during 1990, the Caliente lake completely dried out and deposits of liquid sulphur were observed (Oppenheimer 1992). In the following ten years, the restored lake water has shown strong compositional variations (Martinez *et al.* 2000). The hydrothermal activity at the Active Crater has progressively increased since 2001 (Vaselli *et al.* 2003a), and from 2006 to 2008 several phreatic eruptions took place (Fernandez *et al.* 2007; Duarte & Fernandez 2007). Presently, strong fumaroles discharge hot (up to $322 \text{ }^{\circ}\text{C}$ in November 2008) magmatic-related fluids from the lake bottom and the E-NE inner flank.

2.2 Irazú volcano

Irazú volcano, a composite stratovolcano (3432 m a.s.l.), lies about 25 km E of San José (Fig. 1), and has a

very complex structure, being constituted by several summit craters and a large number of parasitic cones on its flanks. Its edifice, having a predominant calc-alkaline basaltic-andesite and andesitic composition, developed in the early Holocene (Krushensky 1972; Clark 1993). The volcano summit consists of an E-W oriented ridge, likely produced by gravitational collapses (Pavanelli *et al.* 2004), formed by several independent vents. The historical eruptions are mainly characterized by strombolian and vulcanian activity from the westernmost crater (Main Crater). The last event, in 1963-65 (Murata *et al.* 1966), produced more than 40 lahars, some of which reached the city of Cartago. In 1994, part of the northern flank collapsed generating a large debris-flow that flowed N for several tens of km (Barboza *et al.* 1994) likely triggered by a phreatic explosion (e.g., Alvarado *et al.* 2006, Clark *et al.* 2006). Presently, fumarole emissions from the northern outer flank reach 80 °C (March 2008), likely related to a well-developed hydrothermal system, whereas Main Crater, which has a diameter of 690 m and a depth of 150 m, hosts a permanent lake where bubbling gas emissions occur mainly in the NE part of the crater.

2.3 Bosque Alegre caldera (Congo volcano)

The Congo volcanic complex, located 10 km N of Poás volcano, formed about 3000 years ago (Melson *et al.* 1988). It likely represents the remnant of a former twin volcano, whose southern cone (Cerro Congo) is the only one still preserved (Bergoeing & Brenes 1978). Little information is available about the geological and structural features of the Bosque Alegre caldera, located in the NE flank of the Congo volcano (Dondoli *et al.* 1968; Saenz 1971). Three lakes are situated within the Bosque Alegre caldera (740 m a.s.l.): Hule, the largest one, Congo and Bosque Alegre. Advances in accelerator mass spectrometry (AMS) radiocarbon dating of fragments of leaves from terrestrial plants and wood recovered in sediments from the Hule lake indicate that the Hule maar formed between 2950 and 2471 calibrated (cal.) years B.P. and that at least two intra-maar eruptions occurred, about 1740 and 670 cal. years B.P. (Horn 2001). Two events of lake overturn likely occurred prior to 1989 and in 1997, respectively (OVSICORI-UNA 1996; Gocke 1996-1997; Haberyan & Horn 1999). As also reported by nearby settlers, these phenomena were marked by sudden colour changes and death of fishes and birds.

2.4 Tenorio volcano

Tenorio (1916 m a.s.l.) is a complex basaltic-andesite stratovolcano pertaining to the Cordillera de Guanacaste. The edifice consists of two twin craters aligned along the NNW-ESE regional trend. The Tenorio volcanic complex lies on a Pre-Pleistocene basaltic-andesite basement covered by eruptive sequences consisting of alternate lava and pyroclastic deposits (Herrera

Cabezas 1990). No historic volcanic activity has been reported from this volcano, although a legend known by the local population tells about an eruption (never documented) in 1816 (Alvarado 1993). Fumarolic activity and thermal discharges are relatively widespread in the outer NE flank of the volcano (Vaselli *et al.* 2003b). Exploratory geothermal wells have produced no significant results (Moya *et al.* 2006; Moya & DiPippo 2007). Tenorio lake is located in a remote area close to the volcano summit.

2.5 Rincón de la Vieja volcano

Rincón de la Vieja (1916 m a.s.l.), located at about 25 km NE of the town of Liberia (Fig. 1), is a basaltic-andesite stratovolcano (Kempton *et al.* 1996) whose summit consists of seven NW-oriented nested craters. Early volcanism was dominated by lava flows, whereas plinian activity had increased with time (Barquero & Segura 1983). The last eruption with a significant juvenile magmatic component took place 3500 years B.P. (Alvarado *et al.* 1992). Frequent phreato-magmatic and phreatic eruptions have occurred since 1851 (Boudon *et al.* 1996). Several thermal springs presently discharge from the outer flanks of the volcanic edifice (Kempton & Rowe 2000; Tassi *et al.* 2005). The active crater, which hosts the highly acidic Rincón Crater lake, is characterized by intense fumarolic emission (Tassi *et al.* 2005) seeping out from the lake and the inner crater walls. A cold crater lake is located within the Santa Maria volcano, and a permanent freshwater lake ("Laguna") is near the active crater (Baker 1987).

3. MATERIAL AND METHODS.

3.1 Bathymetry and sampling techniques

The collection of water and dissolved gas samples along the vertical profiles of the Hule, Irazú, Botos (February 2001) and Caliente (February 2001 and 2007) lakes was carried out with a tube (6 mm in diameter) of Rilsan, a plastic material impermeable to both water and gas. Once one end of the Rilsan tube was lowered to the chosen depth, water was pumped up to the surface by means of a 150 mL glass syringe, connected to the other end of the tube *via* a three way valve, and transferred into the storage containers (Tassi *et al.* 2004b; 2009). All water samples were collected after the displacement of a water volume double than that of the inner volume of the tube (about 0.03 dm³ m⁻¹). The free gas phase released from water during the suction operated by the syringe was stored into 100 mL glass flasks tapped with teflon valves. To perform water and gas sampling along the vertical profiles, the point above the deepest spot of each lake was reached by using a 2.30 m long inflatable raft. A wooden raft was assembled and used under a rubber raft covered with two folded plastic tarps to minimize the contact with the highly acidic water of the Caliente lake. The sampling profiles are described in table 1.

Tab. 1. Temperatures (°C), pH, chemical composition and $\delta^{18}\text{O}$ and $\delta^2\text{H}$ values (expressed as ‰ V-SMOW) of the crater lakes of Costa Rica. Ion contents are in mg L⁻¹, n.a.: not analyzed.

lake	depth (m)	volcano	date	T °C	pH	Ca ²⁺	Mg ²⁺	Na ⁺	K ⁺	HCO ₃ ⁻	SO ₄ ²⁻	Cl ⁻	F ⁻	Br ⁻	NO ₃ ⁻	B	Li ⁺	NH ₄ ⁺	Fe ³⁺	Mn ²⁺	Al ³⁺	SiO ₂	CO ₂ calc	TDS	$\delta^{18}\text{O}$	$\delta^2\text{H}$
Hule	1	Congo	Feb-01	26.5	6.45	6.7	2.05	3.06	0.87	43.3	1.45	1.40	0.04	0.02	1.34	0.29	<0.02	0.41	2.51	0.20	<0.01	24.0	58	87.6	-4.6	-22.1
Hule	10	Congo	Feb-01	25.3	6.41	6.8	1.94	3.11	0.88	46.4	1.05	1.40	0.05	0.02	0.65	0.27	<0.02	0.89	3.53	0.30	<0.01	28.0	56	95.2	-4.7	-26.5
Hule	20	Congo	Feb-01	26.0	6.24	7.0	1.87	3.24	0.88	44.5	1.30	1.40	0.05	0.02	1.15	0.22	<0.02	0.51	3.47	0.30	<0.01	29.0	87	94.9	-4.6	-20.5
Congo	1	Congo	Feb-01	20.0	5.40	0.9	0.25	0.82	0.16	4.90	0.60	1.20	0.02	0.01	<0.05	0.05	<0.02	0.13	0.01	<0.01	5.00	49	14.3	-3.4	-19.9	
Irazú	1	Irazú	Feb-01	16.0	6.59	204	235	149	48.2	305	1275	230	6.20	0.55	0.80	1.13	0.04	0.08	1.90	1.28	<0.01	178	73	2636	-7.8	-67.8
Irazú	10	Irazú	Feb-01	15.0	6.60	198	225	144	47.4	303	1275	213	5.90	0.55	0.95	1.55	0.04	0.24	0.65	1.26	<0.01	182	414	2598	-8.0	-62.3
Irazú	20	Irazú	Feb-01	15.0	6.55	204	230	149	48.7	299	1260	223	6.10	0.55	0.95	1.67	0.05	0.18	1.96	1.27	<0.01	178	466	2603	-7.8	-61.5
Irazú	30	Irazú	Feb-01	14.5	6.30	201	230	155	47.6	284	1275	225	6.10	0.55	0.90	1.56	0.04	0.20	1.80	1.27	<0.01	177	513	2607	-8.0	-66.9
Tenorio	1	Poás	Feb-01	16.0	5.84	0.6	0.36	2.03	0.16	5.49	1.20	3.25	0.02	0.01	0.23	0.05	<0.02	0.06	0.73	0.02	<0.01	10.0	99	24.2	-2.3	-7.0
Botos	1	Poás	Feb-01	16.8	5.03	2.9	0.59	0.59	0.16	0.61	12.4	0.85	0.09	0.01	0.05	0.52	<0.02	0.24	0.09	<0.01	0.10	14.0	154	33.2	-6.2	-40.6
Botos	8	Poás	Feb-01	12.5	4.40	3.1	0.72	0.67	0.16	0.00	9.75	3.90	0.25	0.02	0.15	0.22	<0.02	0.08	0.06	0.01	0.20	26.0	187	45.2	-5.7	-42.4
Caliente	1	Poás	Feb '08	43.0	0.60	800	644	770	320	0.00	60,398	21,814	1334	43.0	19.00	20.4	0.20	59.0	n.a.	n.a.	n.a.	n.a.	86,222	n.a.	n.a.	n.a.
Caliente	5	Poás	Feb '08	35.0	0.54	812	675	783	333	0.00	63,796	23,240	1390	51.0	49.00	22.1	0.21	76.0	n.a.	n.a.	n.a.	n.a.	91,227	n.a.	n.a.	n.a.
Caliente	10	Poás	Feb '08	38.0	0.46	780	625	806	341	0.00	63,859	23,382	1410	64.0	13.00	21.8	0.20	76.0	n.a.	n.a.	n.a.	n.a.	91,378	n.a.	n.a.	n.a.
Caliente	20	Poás	Feb '08	39.5	0.38	772	627	783	337	0.00	62,678	24,271	1419	57.0	13.00	22.1	0.20	67.0	n.a.	n.a.	n.a.	n.a.	91,046	n.a.	n.a.	n.a.
Caliente	26	Poás	Feb '08	40.0	0.35	794	681	781	321	0.00	63,253	23,088	1487	53.0	10.00	21.3	0.15	59.0	n.a.	n.a.	n.a.	n.a.	90,548	n.a.	n.a.	n.a.
Rin. Crater	0	R. de la Vieja	Feb-98	47.0	1.12	1567	2037	609	292	0.00	12,500	18,800	1400	80.0	10.0	22.0	0.78	254	3720	70.0	n.a.	n.a.	41,362	n.a.	n.a.	n.a.
Rin. Crater	0	R. de la Vieja	Feb-99	31.2	1.16	1300	595	535	130	0.00	9000	8800	370	25.0	22.5	8.64	0.26	52.2	1620	44.2	1060	251	23,814	4.9	1.1	n.a.
Rin. Crater	0	R. de la Vieja	Feb-00	37.0	1.11	1233	500	397	123	0.00	9400	9500	450	21.5	3.95	4.88	0.18	45.5	925	25.0	650	n.a.	23,278	n.a.	n.a.	n.a.
Rin. Crater	0	R. de la Vieja	Feb-01	44.0	0.70	498	433	448	123	0.00	15,750	15,900	660	37.5	11.5	13.8	0.31	125	1024	26.5	940	262	36,272	9.9	2.1	n.a.
Rin. Crater	0	R. de la Vieja	Mar-01	37.0	0.71	1835	755	548	136	0.00	19,250	20,500	1625	37.5	<0.05	15.9	n.a.	183	n.a.	n.a.	n.a.	n.a.	44,905	n.a.	n.a.	6.1
Rin. Crater	0	R. de la Vieja	Jul-01	39.0	0.65	1585	895	900	459	0.00	37,500	29,000	2950	62.5	<0.05	13.1	n.a.	250	n.a.	n.a.	n.a.	n.a.	73,615	5.3	6.1	n.a.
Rin. Crater	0	R. de la Vieja	Mar-02	39.0	0.65	1495	556	947	610	0.00	37,500	29,000	2950	62.5	<0.05	13.1	n.a.	250	n.a.	n.a.	n.a.	n.a.	73,383	5.4	-6.4	n.a.
Rin. Crater	0	R. de la Vieja	Apr-02	37.0	0.30	887	2587	2837	1792	0.00	80,000	63,000	2350	140	35.00	43.0	1.70	649	n.a.	n.a.	n.a.	n.a.	154,322	7.9	8.3	n.a.
Rin. Crater	0	R. de la Vieja	Jul-04	28.5	n.a.	1112	781	912	587	0.00	60,000	21,000	2050	81.0	21.10	15.0	0.70	n.a.	n.a.	n.a.	n.a.	n.a.	173,733	n.a.	n.a.	n.a.
Rin. Crater	0	R. de la Vieja	May-03	35.0	n.a.	912	1169	1220	742	0.00	97,500	55,000	2780	155	11.00	29.0	0.90	411	n.a.	n.a.	n.a.	n.a.	159,930	n.a.	n.a.	n.a.
Rin. Crater	0	R. de la Vieja	Mar-04	32.0	n.a.	1087	931	1037	702	0.00	115,000	51,000	3750	195	13.10	17.0	0.70	n.a.	n.a.	n.a.	n.a.	n.a.	173,733	n.a.	n.a.	n.a.
Rin. Crater	0	R. de la Vieja	Oct-04	49.0	n.a.	1026	850	890	610	0.00	70,000	23,500	2250	89.0	20.50	16.0	0.60	n.a.	n.a.	n.a.	n.a.	n.a.	99,252	9.9	5.4	n.a.
Caliente	0	Poás	Feb-98	37.0	0.89	1130	534	498	69.0	0.00	8250	8000	480	22.5	5.00	6.80	0.17	62.0	n.a.	n.a.	n.a.	n.a.	86,560	n.a.	n.a.	n.a.
Caliente	0	Poás	Feb-99	31.0	0.38	1040	483	508	90.0	0.00	9750	12,800	700	27.5	25.0	3.11	0.15	138	1008	28.4	450	220	27,271	1.5	-22.9	n.a.
Caliente	0	Poás	Nov-99	25.0	0.75	745	285	600	59.0	0.00	5500	4750	325	20.0	19.0	2.10	0.06	24.5	457	12.6	875	195	12,994	-2.0	-26.3	n.a.
Caliente	0	Poás	Feb-00	46.0	1.06	1110	490	417	102	0.00	9200	9400	700	30.0	5.20	6.19	0.11	52.7	875	22.8	390	n.a.	22,801	3.4	-19.6	n.a.
Caliente	0	Poás	May-00	36.0	1.01	1086	930	502	60.0	0.00	12,000	5750	500	15.0	4.00	7.98	n.a.	92.9	n.a.	n.a.	n.a.	n.a.	20,948	n.a.	n.a.	n.a.
Caliente	0	Poás	Jun-00	38.0	0.61	904	482	474	94.5	0.00	11,250	8375	625	28.0	3.50	7.46	n.a.	108	n.a.	n.a.	n.a.	n.a.	22,352	0.8	-21.6	n.a.
Caliente	0	Poás	Jul-00	33.0	0.87	966	540	476	66.0	0.00	15,750	7500	550	17.5	4.50	5.52	n.a.	40.0	n.a.	n.a.	n.a.	n.a.	25,916	0.7	-17.9	n.a.
Caliente	0	Poás	Nov-01	17.0	2.14	901	430	315	86.5	0.00	7750	5750	360	15.0	4.00	4.45	n.a.	104	n.a.	n.a.	n.a.	n.a.	15,720	-0.7	-27.0	n.a.
Caliente	0	Poás	Feb-02	27.0	1.98	644	219	225	38.0	0.00	14,000	18,000	400	51.0	2.50	2.70	0.10	140	n.a.	n.a.	n.a.	n.a.	33,722	n.a.	n.a.	n.a.
Caliente	0	Poás	Apr-02	25.0	1.91	912	350	320	58.0	0.00	8500	4000	225	25.0	5.00	4.10	0.20	47.7	n.a.	n.a.	n.a.	n.a.	14,447	n.a.	n.a.	n.a.
Caliente	0	Poás	May-04	25.0	0.86	869	300	288	60.0	0.00	12,000	8375	450	31.5	9.50	3.70	0.20	n.a.	n.a.	n.a.	n.a.	n.a.	22,387	n.a.	n.a.	n.a.
Caliente	1	Poás	Feb-01	30.8	0.83	496	461	438	78.0	0.00	9250	10,900	730	22.0	2.70	2.74	0.19	47.6	830	26.4	1020	272	24,576	1.3	-22.9	n.a.
Caliente	10	Poás	Feb-01	32.0	0.83	518	485	452	79.5	0.00	10,500	11,600	680	26.5	3.00	7.20	0.19	45.0	845	27.0	970	268	26,506	1.1	-26.1	n.a.
Caliente	20	Poás	Feb-01	32.0	0.83	520	501	466	82.5	0.00	11,000	11,900	720	27.5	<0.05	8.92	0.20	64.0	857	26.0	920	262	27,355	1.5	-26.9	n.a.
Caliente	30	Poás	Feb-01	32.0	0.83	524	504	470	82.0	0.00	9750	11,300	740	23.5	<0.05	7.08	0.20	50.8	883	27.8	860	260	25,482	1.4	-26.9	n.a.
Caliente	41	Poás	Feb-01	32.5	0.83	508	493	452	82.8	0.00	9750	11,400	890	25.5	<0.05	6.62	0.20	41.1	865	27.6	1190	304	26,036	1.4	-22.6	n.a.

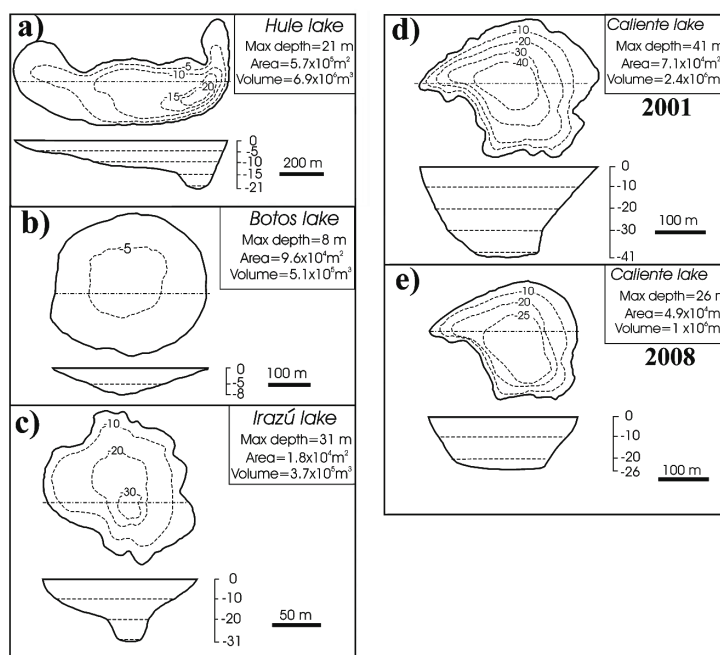


Fig. 2. Reconstruction of the shape and the morphology of the (a) Hule, (b) Botos, (c) Irazú and (d-e) Caliente lakes.

At the Tenorio and Congo lakes, where the maximum depths are 2 and 14.9 m, respectively (Horn & Haberyan 1993), water samples were taken only from the lake surface. Aside from direct sampling performed in February 2001 (Tassi *et al.* 2003a), the high flux of acidic gases released from the Rincón Crater lake has prevented close access and vertical sampling in the lake. Thus, lake waters were collected from the crater rim with a fishing pole bearing a plastic bottle.

The bathymetric profiles of the Hule, Caliente, Botos and Irazú lakes were carried out with a Humminbird Legend 1005 portable Eco-sonar. The morphologic images were produced with the Surfer 6.0 software package.

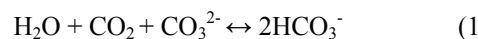
3.2. Analytical methods

Temperature and pH values and HCO_3^- concentrations (acidimetric titration) were measured in the field. Water samples were analyzed for anions (Cl^- , SO_4^{2-} , NO_3^- , Br^- and F^-) by ion-chromatography (Dionex DX-100), whereas cations (Na^+ , K^+ , Ca^{2+} , Mg^{2+} , Fe^{3+} , Mn^{2+} and Li^+) were determined by AAS (Perkin-Elmer AAnalyst 100). SiO_2 , Al^{3+} , NH_4^+ and Boron were analyzed by molecular spectrophotometry. The analytical errors for IC, AAS and colorimetry were 5, 3 and $\leq 10\%$, respectively.

The $^{18}\text{O}/^{16}\text{O}$ and $2\text{H}/1\text{H}$ isotopic ratios (expressed as $\delta^{18}\text{O}$ and $\delta\text{D}\text{‰}$ V-SMOW, respectively) were determined with a Finnigan Delta Plus XL mass spectrometer. Water samples were equilibrated with CO_2 for the analysis of the $\delta^{18}\text{O}$ values (Epstein & Mayeda 1953). The δD values were measured on H_2 obtained after the reaction of 10 μL of water with metallic zinc at 500 $^\circ\text{C}$, following the analytical method proposed by

Coleman *et al.* (1982). EEZ-3 and EEZ-4, calibrated vs V-SMOW and SLAP reference standards, were used as external standards. The experimental error was ± 0.1 and $\pm 1\%$ for the $\delta^{18}\text{O}$ and δD values, respectively.

The composition of the inorganic compounds in the dissolved gas phase (CO_2 , N_2 , O_2 , Ar, Ne, H_2 and He) was determined with a gas-chromatograph (Shimadzu 15a) equipped with a Thermal Conductivity Detector (TCD). Methane and CO (after conversion to CH_4 at 400 $^\circ\text{C}$ using a Shimadzu MTN-1 methanizer), were analyzed with a Shimadzu 14a gas-chromatograph equipped with a Flame Ionization Detector (FID) (Tassi *et al.* 2004a). Dissolved CO_2 concentrations (Tab. 1) were calculated from the total carbonate concentrations, measured by automatic titration (Titroprocessor Metrohm 636) with 0.5M HCl, in water samples collected in 100 mL glass bottles containing 10 mL of a 1M Na_2CO_3 solution, according to the following reaction:



The $^{13}\text{C}/^{12}\text{C}$ ratio (expressed as $\delta^{13}\text{C}\text{‰}$ V-PDB) of CO_2 collected from the Caliente lake was determined with a Finnigan Delta S mass spectrometer after standard extraction and purification procedures of the gas mixtures (Evans *et al.* 1998; Vaselli *et al.* 2006). Internal (Carrara and San Vincenzo marbles) and international (NBS18 and NBS19) standards were used for estimation of external precision. Analytical error is $\pm 0.05\%$.

4. RESULTS

4.1. Bathymetry and shape of crater lakes

Bathymetric maps for the Hule, Botos, Caliente and Irazú lakes are shown in figure 2a-d.

Hule lake is characterized by an uncommon half-moon shape (Fig. 2a), because of the presence of a secondary cone near the northern shore, and which separates Hule and Congo lakes. Hule lake has a maximum depth of 21 m close to the SW shore, and a volume of $6.9 \times 10^6 \text{ m}^3$. Botos and Irazú lakes have a regular circular shape (Fig. 2b,c) with mean diameters of about 350 and 150 m, maximum depths of 8 and 31 m and volumes of 5.1×10^5 and $3.7 \times 10^5 \text{ m}^3$, respectively. In February 2001 the Caliente lake had a mean diameter of 300 m, a maximum depth of 41 m and a volume of $2.4 \times 10^6 \text{ m}^3$ (Fig. 2d). However, the water level has strongly changed in time. In March 2008, the lake diameter, maximum depth and volume were 250 m, 26 m and $1 \times 10^6 \text{ m}^3$, respectively (Fig. 2e). The main morphological features of the Costa Rica crater lakes can be summarized by a dimensionless parameter, known as "depth-ratio" (Carpenter 1983), which is the ratio between the average depth and the maximum depth, the former being equal to the ratio between the volume and the surface area of the lake. This parameter may predict if a lake can be affected by stratification, e.g. the higher the depth-ratios, the higher the probability of stratification (Hutchinson 1957). The depth-ratios of Botos, Hule, and Irazú lakes are 0.66, 0.57 and 0.65. According to the heuristic classification of the shapes of lake basins proposed by Carpenter (1983), these values are consistent with the so-called ellipsoid shape (whose typical values are comprised between 0.5 and 0.66), which can be considered a common feature for lake basins in volcanic environments. Conversely, Caliente lake in 2001 and 2007 shows depth-ratio values of 0.82 and 0.78, respectively, approximating a steep-sided frustum model (corresponding to steep sides and flat bottom; Lehman 1975), typical of lakes of recent formation and affected by elevated sedimentation rates.

4.2. Water chemical and isotopic composition

Temperature and pH values, $\delta^{18}\text{O}$ and δD ratios and chemical composition of the crater lakes are listed in table 1. Rincón Crater lake was characterized by a relatively high temperature (up to $49 \text{ }^\circ\text{C}$), caused by the conspicuous inputs from the lake bottom of hot fluids related to the strong activity characterizing the Rincón de la Vieja magmatic-hydrothermal system. Phreatic eruptions have indeed frequently occurred between February 1998 and October 2004, causing strong changes of water level and chemistry (Tassi *et al.* 2005 and references therein). Relatively high temperatures (up to $46 \text{ }^\circ\text{C}$) were also measured at the Caliente lake, likely because of the presence of numerous sub-lacustrine fumaroles fed by a hydrothermal system, permanently affected by inputs of high-temperature volatiles (Vaselli *et al.* 2003a) released from a shallow ($\approx 500 \text{ m}$) magmatic body (Rowe *et al.* 1995). The unusual low temperature ($17 \text{ }^\circ\text{C}$) measured in November 2001 was likely due to the influence of heavy rains just before water

sampling. Temperatures measured at Caliente along the vertical profile in 2008 (up to $43 \text{ }^\circ\text{C}$) indicated the virtual absence of any significant thermal stratification. In 2001 the lake temperatures was relatively low ($<32.5 \text{ }^\circ\text{C}$) (Vaselli *et al.* 2003a), likely in relation to heavy rains affecting the area, as also testified by the relatively large volume of the lake in that period (Fig. 2e, f). No significant temperature variations were recorded in the water columns of the Hule (between 25.3 and $26.5 \text{ }^\circ\text{C}$) and Irazú lakes (between 14.5 and $16 \text{ }^\circ\text{C}$), whereas the temperature measured at the surface of the Botos lake ($16.8 \text{ }^\circ\text{C}$) was significantly higher than that measured at the lake bottom ($12.5 \text{ }^\circ\text{C}$), and was likely because of sun exposure.

Hule, Botos, Congo and Tenorio lakes had an acidic pH (ranging from 4.44 to 6.45) and relatively low TDS (Total Dissolved Solids) values ($<100 \text{ mg L}^{-1}$). Irazú lake had higher TDS ($\approx 2,600 \text{ mg L}^{-1}$) and pH (up to 6.6) values, whereas the Caliente and Rincón Crater lakes were characterized by extremely high TDS values (up to $91,449$ and $173,733 \text{ mg L}^{-1}$, respectively) and extremely low pH (between 0.30 and 2.14). The amount of dissolved solids is on average significantly higher with respect to other highly acidic crater lakes worldwide, e.g. Ruapehu, New Zealand (Christenson 2000), Kawah Ijen, Indonesia (Delmelle *et al.* 2000), Santa Ana, El Salvador (Bernard *et al.* 2004); Copahue, Argentina (Varekamp *et al.* 2009). The Hule and Congo lakes have displayed a $\text{Ca}(\text{Mg})\text{-HCO}_3^-$ composition, typical of worldwide superficial waters and shallow aquifers, whereas the Tenorio lake had a Na-Cl composition (Fig. 3).

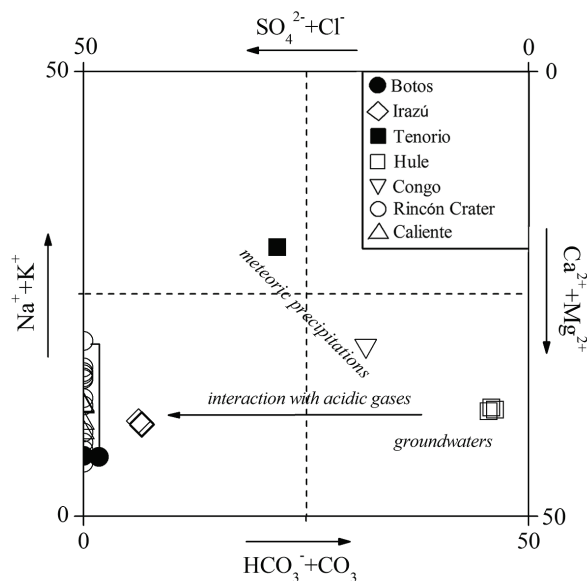


Fig. 3. Langelier-Ludwig (1942) diagram for the water samples from the Hule (open square), Irazú (open diamond), Caliente (open up-triangle), Botos (closed circle), Rincón Crater (open circle), Congo (open down-triangle) and Tenorio (closed square) lakes.

Tab. 2. Chemical composition and $\delta^{13}\text{C}$ in CO_2 (expressed as ‰V-PDB) values of the dissolved gases from crater lakes of Costa Rica. Gas contents are in % by vol.; n.a.: not analyzed.

lake	depth (m)	CO_2	N_2	Ar	O_2	CH_4	Ne	H_2	He	CO	$\delta^{13}\text{C}$
Hule	20	1.713	92.76	2.215	3.315	<0.0001	0.0010	<0.0001	<0.0001	<0.0001	n.a.
Irazú	30	39.47	50.33	0.698	9.441	0.0578	0.0005	0.0011	0.0024	<0.0001	n.a.
Botos	8	0.042	73.06	1.872	25.02	0.0024	0.0009	<0.0001	<0.0001	<0.0001	n.a.
Caliente	1	3.796	76.69	1.954	17.45	<0.0001	0.0009	0.115	<0.0001	<0.0001	n.a.
Caliente	5	6.631	76.13	1.712	15.01	<0.0001	0.0008	0.511	<0.0001	0.0015	-2.9
Caliente	10	12.11	70.54	1.215	15.48	0.0001	0.0008	0.654	<0.0001	0.0025	-3.0
Caliente	20	16.52	68.25	0.874	13.21	0.0001	0.0006	1.145	<0.0001	0.0045	-3.0
Caliente	26	24.95	61.24	0.689	11.88	0.0002	0.0005	1.215	<0.0001	0.0245	-21.0
Caliente	1	2.695	76.72	0.994	18.67	0.0002	0.0006	0.919	<0.0001	0.0035	-3.0
Caliente	20	13.44	67.91	0.954	16.51	0.0001	0.0006	1.181	<0.0001	0.0075	-3.3
Caliente	41	17.91	64.54	0.745	15.78	0.0003	0.0005	1.007	<0.0001	0.0150	-33.6

By considering the extremely low TDS value (24.2 mg L⁻¹), this chemical signature may be ascribed to meteoric precipitation. Botos and Irazú lakes had a similar composition (Mg-SO₄ and Ca-SO₄, respectively), although their TDS values showed a difference of two orders of magnitude (Tab. 1). Caliente and Rincón Crater lakes showed a Ca(Mg)-Cl(SO₄) composition (Fig. 3) that can be regarded as a common feature for crater lakes in active volcanic systems (Casadevall *et al.* 1984; Christenson & Wood 1993; Delmelle & Bernard 1994). Fluoride concentrations were very high at both the Caliente and Rincón Crater lakes (up to 1419 and 3750 mg L⁻¹, respectively), where those of B and NH₄⁺, which are commonly associated to hot-fluid inputs, are enriched in both these lakes (up to 22 and 649 mg L⁻¹, respectively). In contrast, the Irazú, Tenorio, Botos, Congo and Hule lakes had relatively low concentrations of F⁻, B and NH₄⁺ (up to 6.2, 1.67 and 0.89 mg L⁻¹, respectively). Fe³⁺, Mn²⁺ and Al³⁺ concentrations were abundant in the two hyper-acidic lakes (up to 3720, 70 and 1190 mg L⁻¹, respectively), whereas they did not exceed 3.5 (Fe³⁺), 1.3 (Mn²⁺) and 0.2 (Al³⁺) mg L⁻¹ in those lakes characterized by pH ≥ 4.4. Lithium concentrations were relatively low (<1.7 mg L⁻¹) in all the lakes, whereas NO₃⁻ concentrations, present at minor amounts in the Irazú, Tenorio, Botos, Congo and Hule lakes, had extremely strong variations (ranging from <0.05 to 49 mg L⁻¹) in time at Caliente and Rincón Crater lakes. The concentrations of Br⁻ were essentially correlated to those of Cl⁻. Silica concentrations varied from 5 (Congo) to 304 mg L⁻¹ (Caliente lake in 2001 at the depth of 41 m).

The values of $\delta^{18}\text{O}$ and δD values covered a wide range, varying from -8 to 9.9 and -67.8 and 8.3‰ V-SMOW, respectively. The lowest values are those of Irazú, whereas the most positive were related to Rincón Crater. The distribution of the $\delta^{18}\text{O}$ and δD values along the vertical profiles of the Hule, Irazú, Botos and Caliente (2001 data) lakes (Tab. 1) indicated that isotopic stratification can be considered negligible.

4.3. Dissolved gas composition

Chemical composition and $\delta^{13}\text{C}$ -CO₂ values of the dissolved gas samples are shown in table 2. The com-

position of the gas fraction collected by our sampling method depends on the force applied to the syringe, since the composition of the gas phase released from the water differed from that of the whole dissolved gas mixture. Each gas species has indeed a distinct behaviour in response to the stripping process, being characterized by a different solubility. However, these data are useful to detect anomalous concentrations of non-atmospheric originated gas compounds, such as CO₂ and CH₄, as also supported by the results of analogous investigations (e.g., Kusakabe 1996).

Nitrogen was the most abundant gas component in all the analyzed samples, followed by O₂ concentrations that at the Irazú and Caliente lakes were relatively high (9.4 and 15.8% by vol., respectively) even in the deeper water strata. This suggests an efficient water vertical circulation able to promote frequent mixing of shallow and deep waters preventing the formation of anoxic hypolimnion. In contrast, at the bottom of Hule lake the O₂ concentration (3.3% by vol.) was almost one order of magnitude lower than expected on the basis of the O₂/Ar ratios for ASW (20.24 for pure water at 30 °C, respectively). Relatively high dissolved CO₂ concentrations were measured in samples collected from the maximum depths of the Irazú (39.47% by vol.) and Caliente lakes (17.91 and 24.95% by vol. in 2001 and 2007, respectively). Caliente lake water samples have also shown significant concentrations of H₂ and CO (up to 1.18 and 0.015% by vol., respectively). At Botos lake the dissolved gas phase was essentially constituted by atmospheric-related compounds, whereas the Hule lake was marked by a slight enrichment of CO₂ (1.8% by vol.). Methane was measured in significant concentrations only in the Irazú (0.058% by vol.) and Botos lakes (0.0024% by vol.).

The distribution of the $\delta^{13}\text{C}$ -CO₂ values measured in both 2001 and 2008 along the vertical profile of the Caliente lake showed a sharp variation. The samples collected at the lake bottom had indeed an extremely negative carbon signature (down to -33.6 and -21‰ V-PDB, in 2001 and 2008, respectively), whereas the shallower water strata have shown a significant ¹³C-CO₂ enrichment leading to $\delta^{13}\text{C}$ -CO₂ values (between -2.9 and -3.3‰ V-PDB) similar to those measured in the

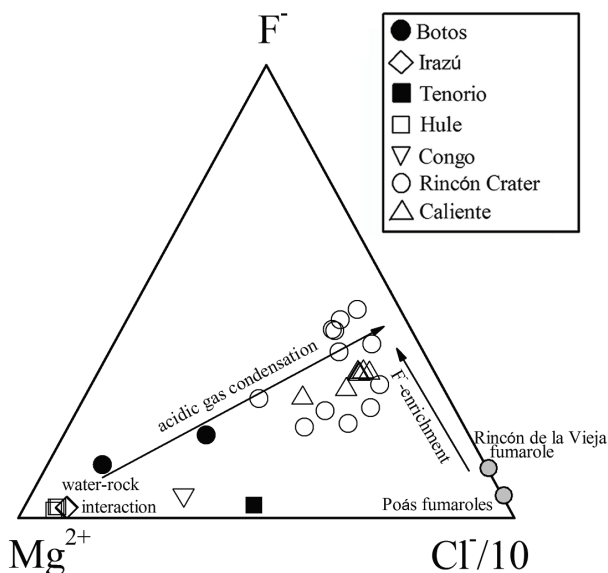


Fig. 4. F^- - $Cl^-/10$ - Mg^{2+} ternary diagram for the water samples from the Hule, Irazú, Caliente, Botos, Rincón Crater, Congo and Tenorio lakes. Symbols as in figure 3.

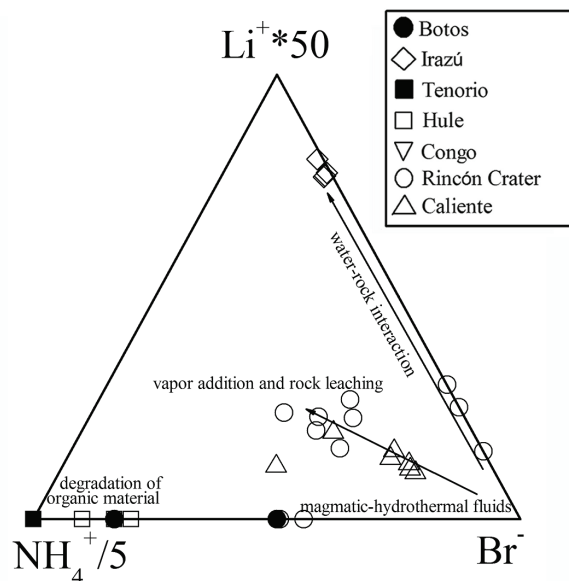


Fig. 5. Br^- - Li^+*100 - $NH_4^+/5$ ternary diagram for the water samples from the Hule, Irazú, Caliente, Botos, Rincón Crater, Congo and Tenorio lakes. Symbols as in figure 3.

fumaroles located N and NE of the lake, i.e. -3 and -6‰ (Vaselli *et al.* 2003a).

5. DISCUSSION

5.1. Processes governing lake-water composition

According to the classification based on the main physical-chemical parameters proposed by Pasternack & Varekamp (1997) and Varekamp *et al.* (2000), the crater lakes of Costa Rica can be grouped, as follows: 1) Rincón Crater and Caliente, which are "high-activity lakes" typical of active volcanic systems (e.g., Ruapheú; Christenson & Wood 1993); 2) Irazú lake, a "low activity lake" related to hydrothermal venting from lake bottom; 3) Hule, Botos, Congo and Tenorio lakes, where no significant degassing activity is present. The chemical features of the Rincón Crater and Caliente lakes are mainly dependent on fluids from the magmatic systems. Therefore, the complex evolution of the chemical-physical features of these two lakes in the last 7 years (Tab. 1) was related to i) the input rate of magmatic-related fluids, regulating lake water evaporation, ii) solute contribution from intense rock leaching and iii) rainfall.

As shown in the F^- - Cl^- - Mg^{2+} ternary diagram (Fig. 4), the Rincón Crater and Caliente lakes were significantly enriched in both Mg^{2+} and F^- with respect to the fumarolic fluids discharged in the same areas (Vaselli *et al.* 2003a; Tassi *et al.* 2005). The Mg^{2+} enrichment was likely due to interaction processes between lake water and host rock. The extremely high F^- concentrations may be caused by dissolution of magmatic-related gas compounds. Crater lakes act indeed as condenser for magmatic volatiles, whose solubility in water depends on pH: at very low pH, HCl

solubility abruptly decreases (Symonds *et al.* 2001) and, consequently, this compound may partly be released to the atmosphere, whereas HF tends to be accumulated in lake water, being completely dissolved. This is also supported by the gas chemistry of Poás (Vaselli *et al.* 2003a) and Rincón de la Vieja (Tassi *et al.* 2005) fumaroles (Fig. 4) that are richer in Cl and poorer in F with the respect to the lake waters. The Irazú and Hule lakes, whose chemical composition is mainly controlled by water-rock interactions, plot close to the Mg^{2+} corner (Fig. 4). The relatively high HCO_3^- and SO_4^{2-} concentrations measured at Irazú (Tab. 1) are related to the CO_2 (H_2S)-rich thermal fluids that are discharged into the lake from the sub-lacustrine vents. Botos, Congo and Tenorio lakes plot at an intermediate position between the fields of hyperacidic and almost neutral lakes (Fig. 4), because they are largely constituted by meteoric water, as also supported by the extremely low TDS values (Tab. 1).

In the NH_4^+ - Li^+ - Br^- ternary diagram (Fig. 5) the different fluid sources feeding the crater lakes can clearly be distinguished. The NH_4^+ enrichment showed by the Tenorio and Congo samples is likely related to degradation of organic material that seems to be the only process active in these lakes, along with the minor contribution of solutes by meteoric waters and water rock interactions. In contrast, Botos lake waters were slightly enriched in Br^- , possibly because of the influence of the Poás fumarolic-induced plume. This may explain the low pH values and the anomalous concentrations of both Cl^- and SO_4^{2-} with respect to those of HCO_3^- (Tab. 1). Irazú was enriched in Li^+ produced by the dominant water-rock interactions. The distribution of the water samples from the Caliente and Rincón Crater lakes (Fig.

5) was depending on the balance between addition of i) HBr and NH_3 from the fumarolic sub-lacustrine discharges and ii) Li^+ from rock leaching.

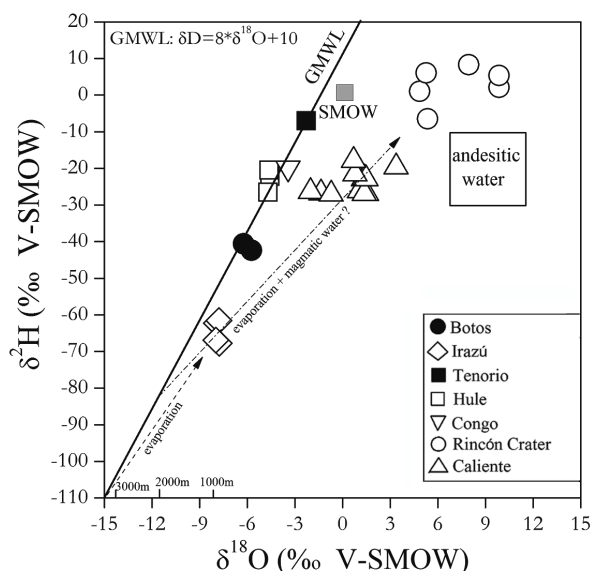


Fig. 6. $\delta^{18}\text{O}$ - δD diagram for the water samples from the Hule, Irazú, Caliente, Botos, Rincón Crater, Congo and Tenorio lakes. The "andesitic water" field is from Taran *et al.* (1989); the Global Meteoric Water Line (GMWL) from Craig (1961). Symbols as in figure 3.

In the $\delta^{18}\text{O}$ vs δD diagram (Fig. 6) the Botos, Tenorio, Hule and Congo lakes plot along the Global Meteoric Water Line (GMWL; Craig 1961). In contrast, the Rincón Crater samples, as well as those of Caliente lake, show a strong positive ^{18}O -shift that may be attributed to evaporation related to the heating of the lake water, mainly caused by hot fluid inputs. As also suggested by Tassi *et al.* (2005), such process has presumably masked any isotopic signature deriving from direct contribution of magmatic waters ("Andesitic water"; Taran *et al.* 1989). The $\delta^{18}\text{O}$ and δD values of the Irazú lake seem to be exclusively related to evaporation of meteoric water, once the effect of the altitude (up to 3300 m a.s.l.) of the recharge area of this lake is considered (Fig. 6).

5.2. Chemical and isotopic features of lakes along the vertical profiles

The chemical and isotopic water compositions of the Irazú, Botos and Hule lakes were basically constant along the vertical profiles (Tab. 1). However, the Irazú and Hule lakes show relative increases with depth of the dissolved CO_2 concentrations determined from both the calculated CO_3^{2-} excess [see reaction (1)] (Tab. 1) and the analysis of the gas phase extracted from water (Tab. 2). The presence of dissolved- CO_2 in the deep strata of the Irazú lake is related to the inputs from the underlying hydrothermal system through the numerous

$\text{CO}_2(\text{H}_2\text{S})$ -rich gas vents that discharge from lake bottom (Tassi *et al.* 2003b) and that were also responsible for the SO_4 -dominated composition. At the Hule lake, the relative small increase of CO_2 with depth could be ascribed to degradation of organic material, in agreement with the relatively high concentrations of N-compounds along the whole vertical profile, typically produced by bacterial activity (Gocke 1996-1997). However, radiocarbon dating of bulk surface sediments and core samples indicates that isotopically dead carbon entered into the lake (Horn 2001), suggesting that CO_2 is released into the lake from a deep source, although at relatively low amounts by considering the low percentage of the measured CO_2 (Tab. 2). This hypothesis is consistent with the lack of CH_4 (Tab. 2), whose production by bacteria is favoured at reducing conditions, such as those dominating the Hule deep water strata. As already suggested by Haberyan & Horn (1999), the relatively low concentrations of the main anions (Tab. 1) seem to exclude direct injection of deepwater of volcanic origin. The shallow Botos lake shows no significant thermal and/or compositional stratification. The distribution of the main chemical compounds from the surface to the bottom of the Caliente lake has been already debated by Vaselli *et al.* (2003a) on the basis of the vertical profile performed in 2001. The main cations have resulted almost constant, whereas F^- showed significant increases. In contrast, the Cl^- and SO_4^{2-} concentrations followed increasing trends from the lake surface to the depth of 20 m (up to 10% of the total abundances) and decreased in the deeper water strata. At the lake bottom both these compounds had concentrations similar to those measured at the surface. The water composition along the vertical profile performed in March 2008 has confirmed the occurrence of this peculiar chemical stratification (Fig. 7a-d). Remarkably, despite the phreatic eruptions occurred from 2006 and the decreasing water level, solute concentrations have not suffered significant variations. These compositional trends may be explained by hypothesizing a complex combination of several processes. Dilution by meteoric precipitations could tentatively be regarded as the main cause for the relatively low Cl^- , SO_4^{2-} and F^- concentrations measured at shallow depth. The summit of Poás volcano is indeed characterized by frequent and heavy rainfall. Nevertheless, the almost constant vertical distribution of the concentrations of the main cations, besides that of $\delta^{18}\text{O}$ and δD values (Tab. 1), seems to exclude that rains can really play a dominant role on the behaviour of the main anions in the shallower strata lake. Evaporation of acidic compounds, like HCl and H_2S , favoured by the extremely low pH and the relatively high temperature of the lake, may better fit with the observed phenomena. Dissolution of HCl, SO_2 , H_2S and HF from the sub-lacustrine fumaroles in lake water may be invoked to explain the progressive enrichment with depth of Cl^- , SO_4^{2-} and F^- . However, only F^- con-

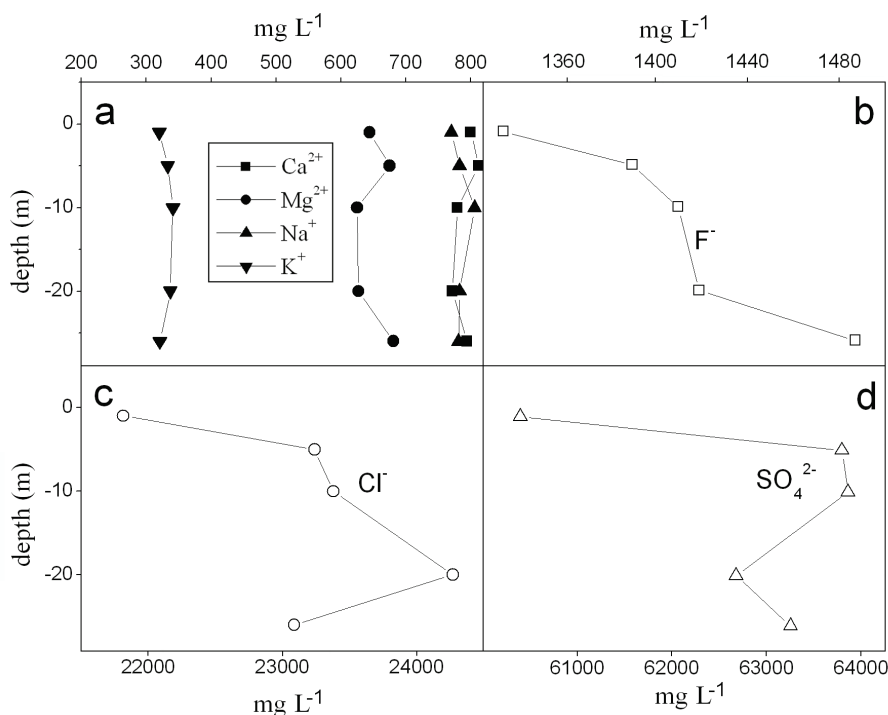


Fig. 7. Vertical profiles of a) Ca²⁺, Mg²⁺, Na⁺ and K⁺, b) F⁻, c) Cl⁻ and d) SO₄²⁻ concentrations for the Caliente lake.

stantly increases with depth. Therefore, at least the behaviour of SO₄²⁻ below the depth of 20 m is likely to be related to the activity of bacteria that are concentrated close to the water-sediment interface. Culture experiments using different media suggested that the bacteria found in the Caliente lake mainly pertain to *Desulfovibrio desulfuricans* and *Desulfovibrio vulgaris* species (Bergamaschi 2002). Such organisms are able to transform SO₄²⁻ to H₂S (Postgate & Campbell 1966). The presence of microbial activity was also supported by the δ¹³C-CO₂ values that from surface to the lake bottom passed from the typical values for mantle-related CO₂ (Rollison 1993) to values characteristic of an organic source (Tab. 2). Although no isotopic data are available for CH₄, the presence of this hydrocarbon in the dissolved gas phase in the deepest portion of Caliente lake (Tab. 2) as originated by biogenic processes cannot be ruled out. Finally, the organic activity is generally considered to be unable to metabolize Cl⁻, although little is known about the characteristics of the bacteria active in such extreme environmental conditions.

6. CONCLUSIONS

The strong fumarolic activity of the Poás and Rincón de La Vieja volcanoes plays an effective control on the chemistry of the two lakes hosted within their main craters. Caliente and Rincón Crater lakes are affected by impressive inputs of hydrothermal/magmatic fluids that have a twofold effect: they add highly soluble compounds and promote intense rock leaching processes. Moreover, the high-flux injection of hot fluids is able to

continuously sustain a convective regime that homogenizes the lake temperature and, at least partly, the chemical composition along the vertical profile. However, biochemical processes at the sediment-water interface appear to significantly affect the chemical and isotopic composition of the bottom waters of Caliente lake. The rapid variations in water level at Caliente lake in the last years is to be ascribed to i) strong evaporation, related to periods of high fumarolic emission, and ii) high-rate sediment deposition from the heavily altered lake border, as also suggested by the permanent flat shape of the lake bottom and iii) the recent phreatic activity. The chemical features of the Irazú lake are dominated by water-rock interaction and CO₂(H₂S)-rich gas addition from the presently-active hydrothermal system. The lack of significant changes with depth of both the water chemistry and temperature may be related to frequent mixing of the lake water, likely triggered by the cold heavy rains commonly affecting this area. The chemical and physical properties of the lakes characterized by extremely low salinity (Botos, Congo and Tenorio) are mainly regulated by meteoric precipitation, although Botos seems to be significantly affected by contamination from the volcanic plume produced by the Active Crater of Poás. At the Hule lake radiocarbon analyses of lake sediments reveals the presence of volcanic CO₂, likely related to the "Poás volcanic alignment" (Malavassi *et al.* 1990), a zone of apparent crustal weakness that begins from the Poás volcano and extends northward through the Hule maar to the Rio Cuarto maar (Alvarado 1990). Episodes of lake turnover are likely frequent in this lake, especially in January, when

the air temperature is relatively cold and the weather is rainy and windy (Haberyan & Horn 1999), as also suggested by the occurrence of fish deaths observed in recent years the local population (Duarte & Fernandez 1996).

The CO₂-rich gas phase detected in the deep waters of Irazú and Caliente lakes is likely from high-flux gas vents discharging from the lake bottoms, although at Caliente the $\delta^{13}\text{C-CO}_2$ values at the maximum depth, as well as the anomalous vertical distribution of the main anions, points to a significant CO₂ contribution from microbial colonies. However, the dimensions of these reservoirs are limited by the continuous gas dispersion through lake surface, which is favoured by the dominant convective regime resulting in frequent mixing of water strata. Therefore, in these systems the hazard related to the possible occurrence of Nyos-type gas eruptions can be considered negligible. Although no vertical chemical and isotopic data at Caliente lake were carried prior to the 2006-2008 phreatic eruptions, it seems that significant variations may have occurred. As a consequence, periodic sampling and analysis of the vertical profiles should be conducted at Caliente lake to minimize the volcanic risk for thousands of tourists that visit the National Park of Poás each year. No measurement along the vertical profile of the Rincón Crater lake are available. However, on the basis of the compositional characteristics of the lake surface it is reasonable to argue that the processes regulating the chemistry of this lake at depth should not differ to those acting at Caliente lake.

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